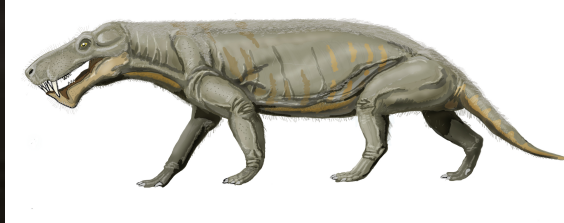


# Ocean Anoxia: The end-Permian Mass Extinction



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NASA ASTROBIOLOGY INSTITUTE



National Science Foundation  
WHERE DISCOVERY BEGINS

PENNSTATE



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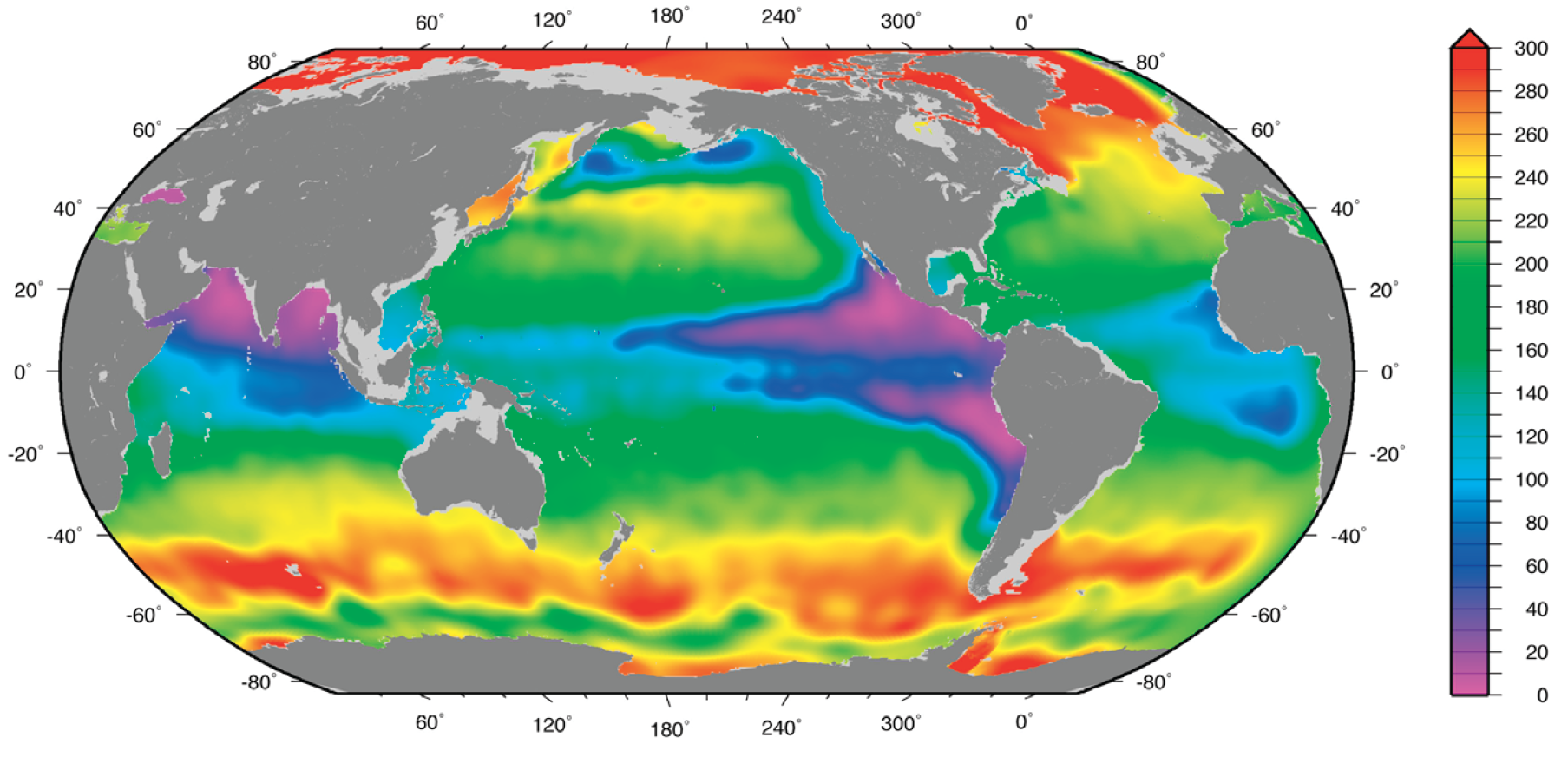
## Ocean Deoxygenation: Past, Present, and Future

PAGES 409–410

oxygen concentrations highlights progress  
scientists have made in understanding oce

periods (385–360 million years ago), wide-  
spread anoxia in shallow continental seas  
coincided with an extended biotic crisis.  
Anoxia was also widespread in both shal-  
low marine and deep ocean environments  
during the Permian-Triassic extinction (~252  
million years ago), marked by the loss of  
approximately 90% of all marine animal

### Oxygen ( $\mu\text{mol/kg}$ ) at 200 m



*Fig. 1. Mean global ocean oxygen concentrations at 200 meters below the surface. Note the extensive regions of low oxygen (oxygen minimum zones) throughout the low-latitude oceans and the subarctic Pacific. Data from the World Ocean Circulation Experiment Global Hydrographic Climatology [Gouretski and Koltermann, 2004].*



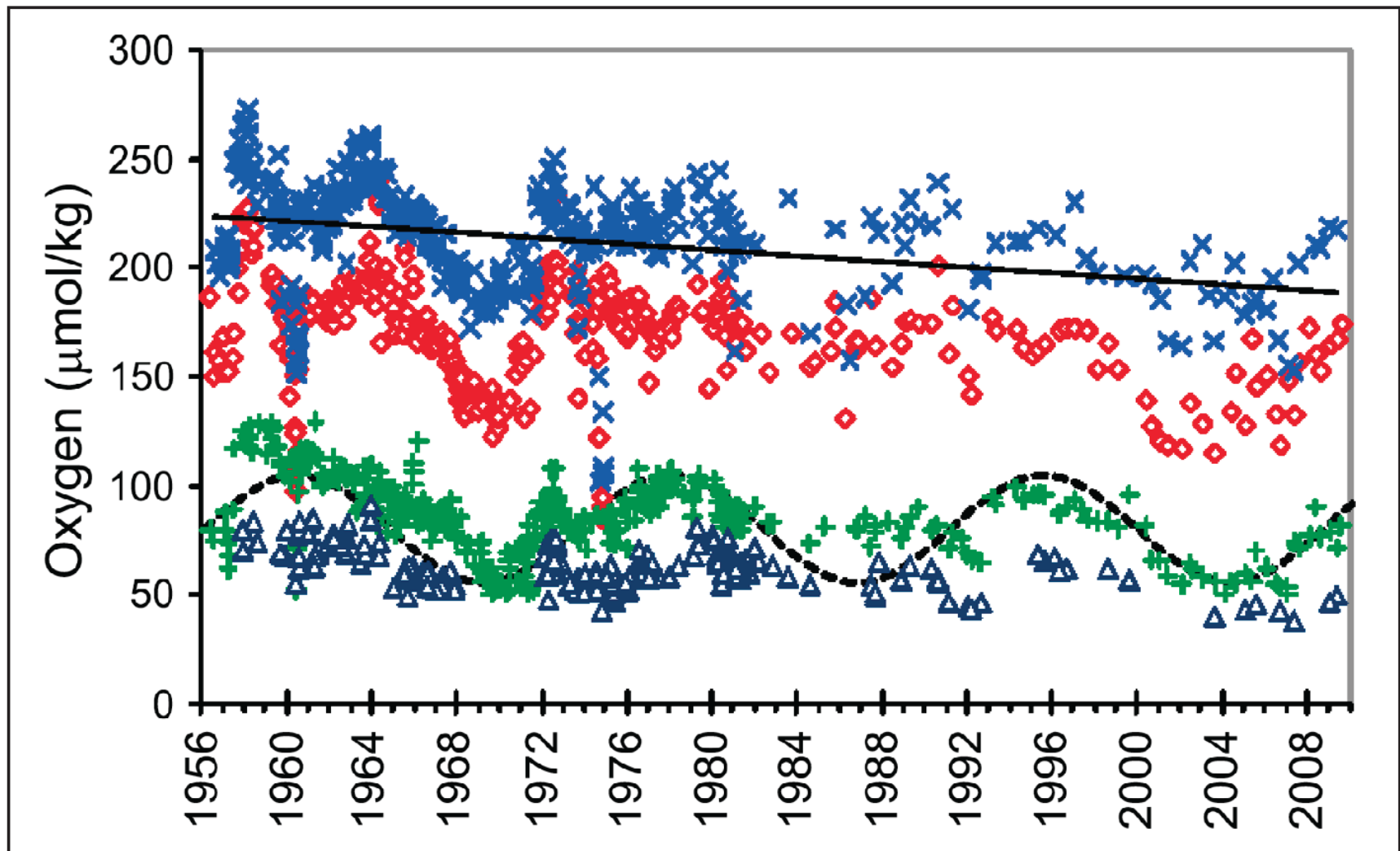
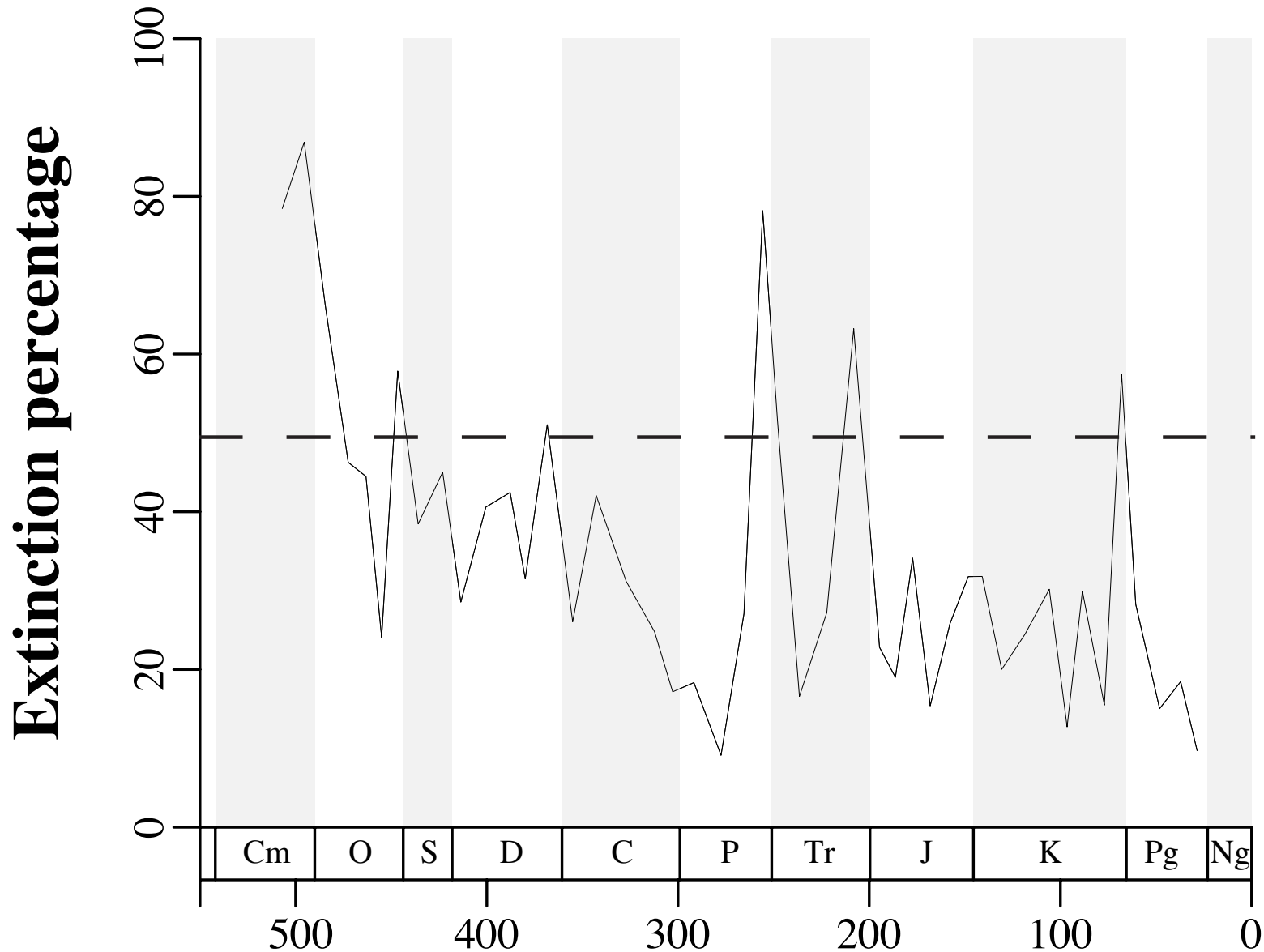


Fig. 2. The decadal trend and oscillations in oxygen concentrations on constant potential density surfaces (isopycnals). Data are shown for isopycnals 26.5 (blue crosses), 26.7 (red diamonds), 26.9 (green pluses), and 27.0 (dark blue triangles) at Ocean Station P ( $50^{\circ}\text{N}$ ,  $145^{\circ}\text{W}$ ). Solid line is the linear regression of the 26.5 isopycnal showing an annual oxygen loss rate of 0.67 micromole per kilogram. The dashed sine wave shows that the oscillation on the 26.9 isopycnal has an amplitude of 50 micromoles per kilogram (centered on 80 micromoles per kilogram) and a period of 18.6 years [Whitney et al., 2007; F. Whitney, personal communication, 2011].

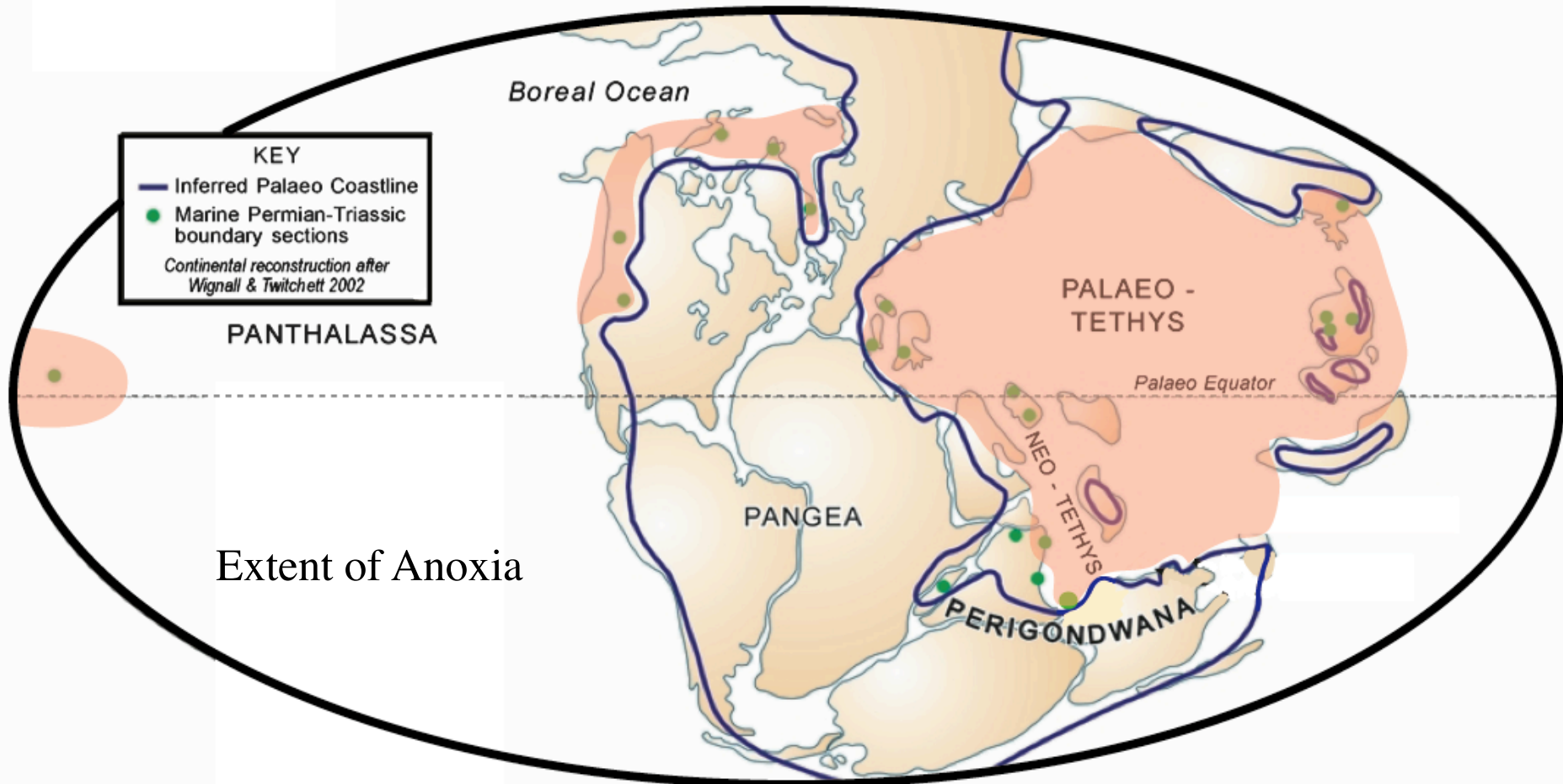




**Time (Ma)**

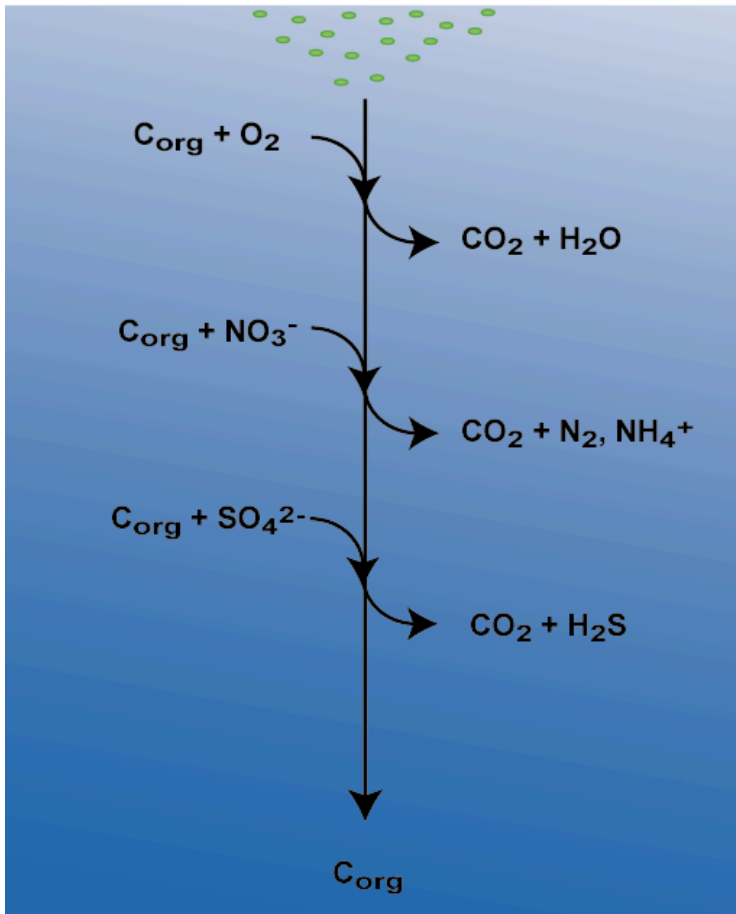
Alroy, pers. comm.

# Permian-Triassic Earth

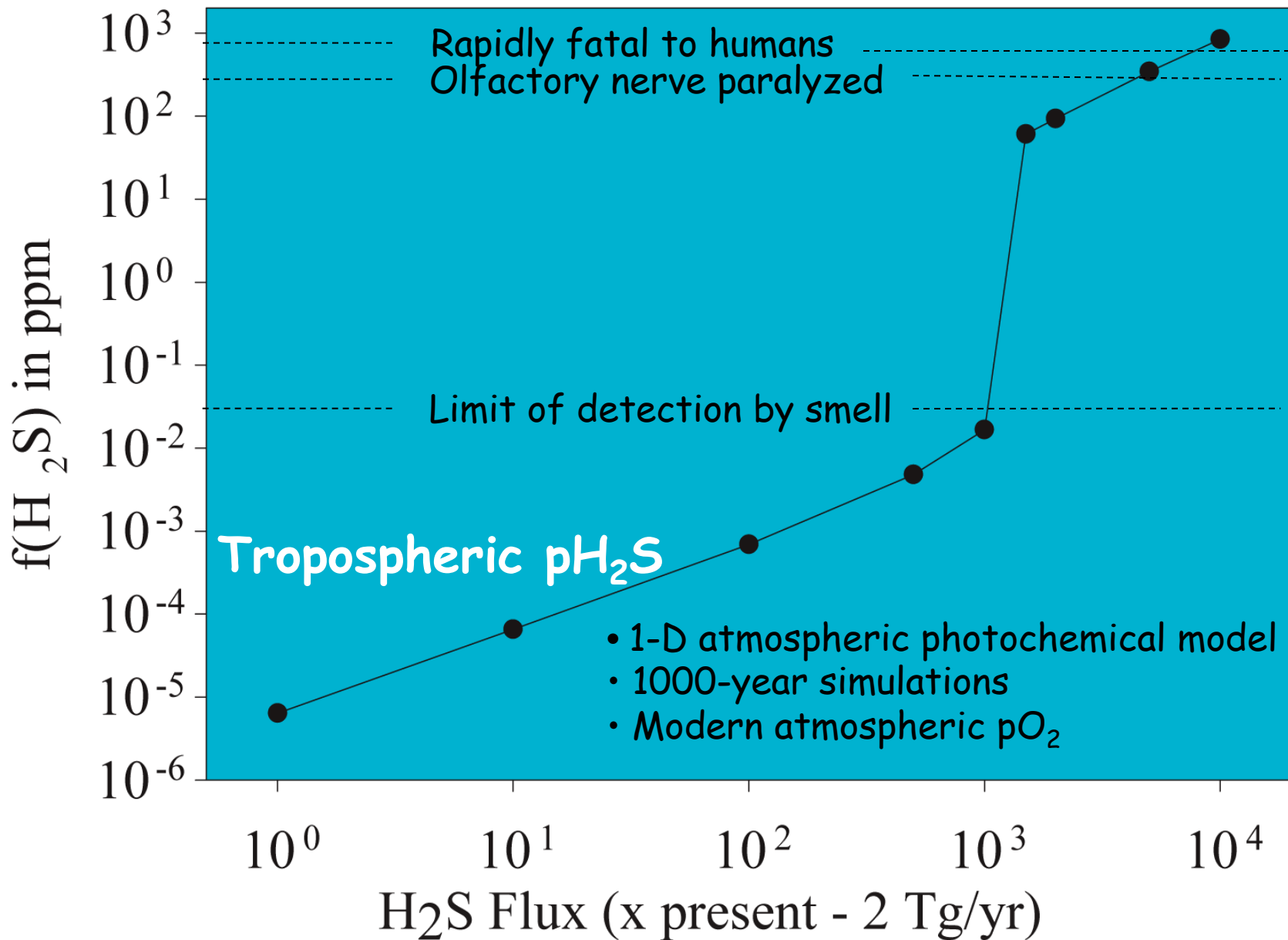


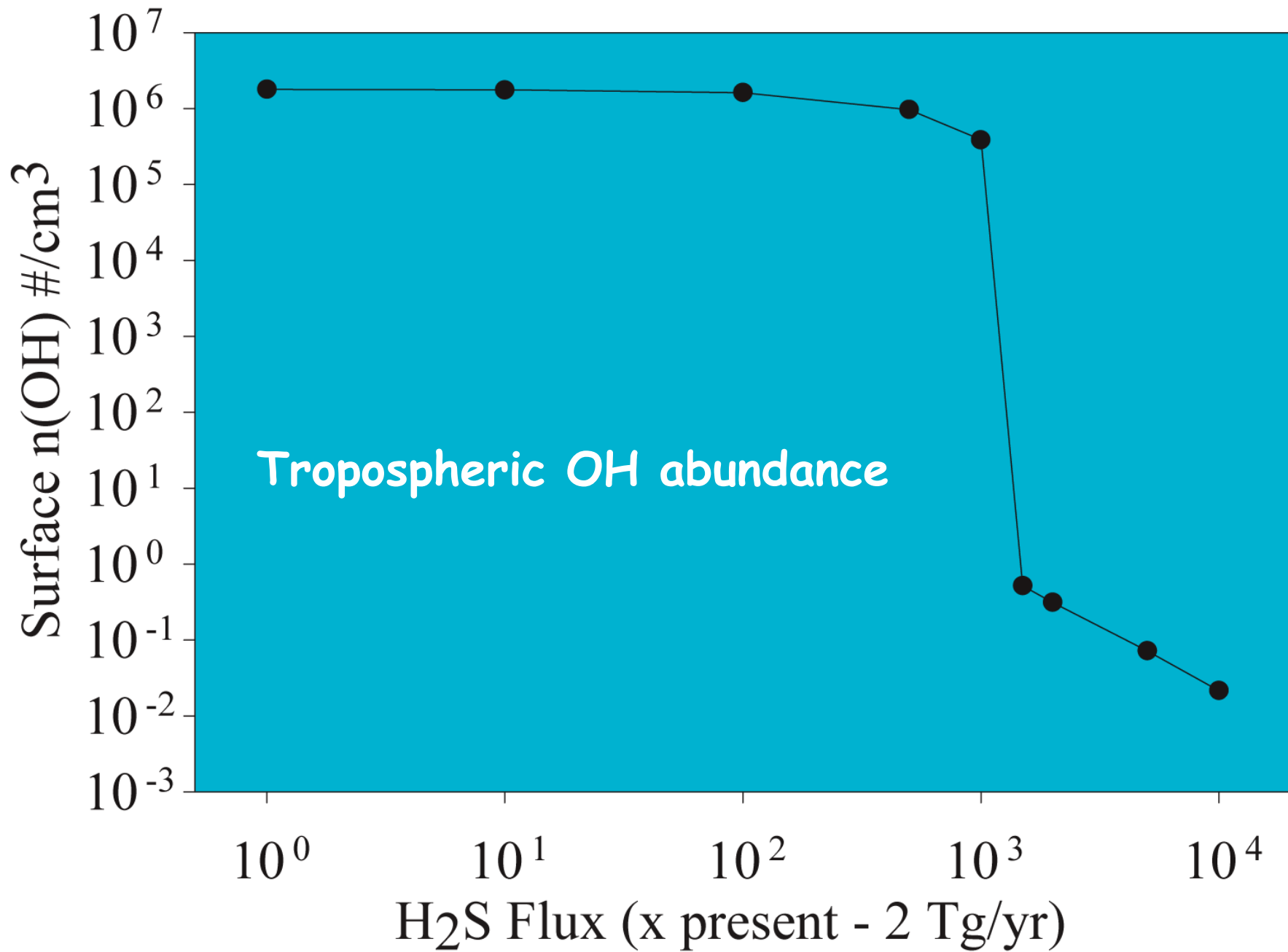
# Photic-Zone Euxinia

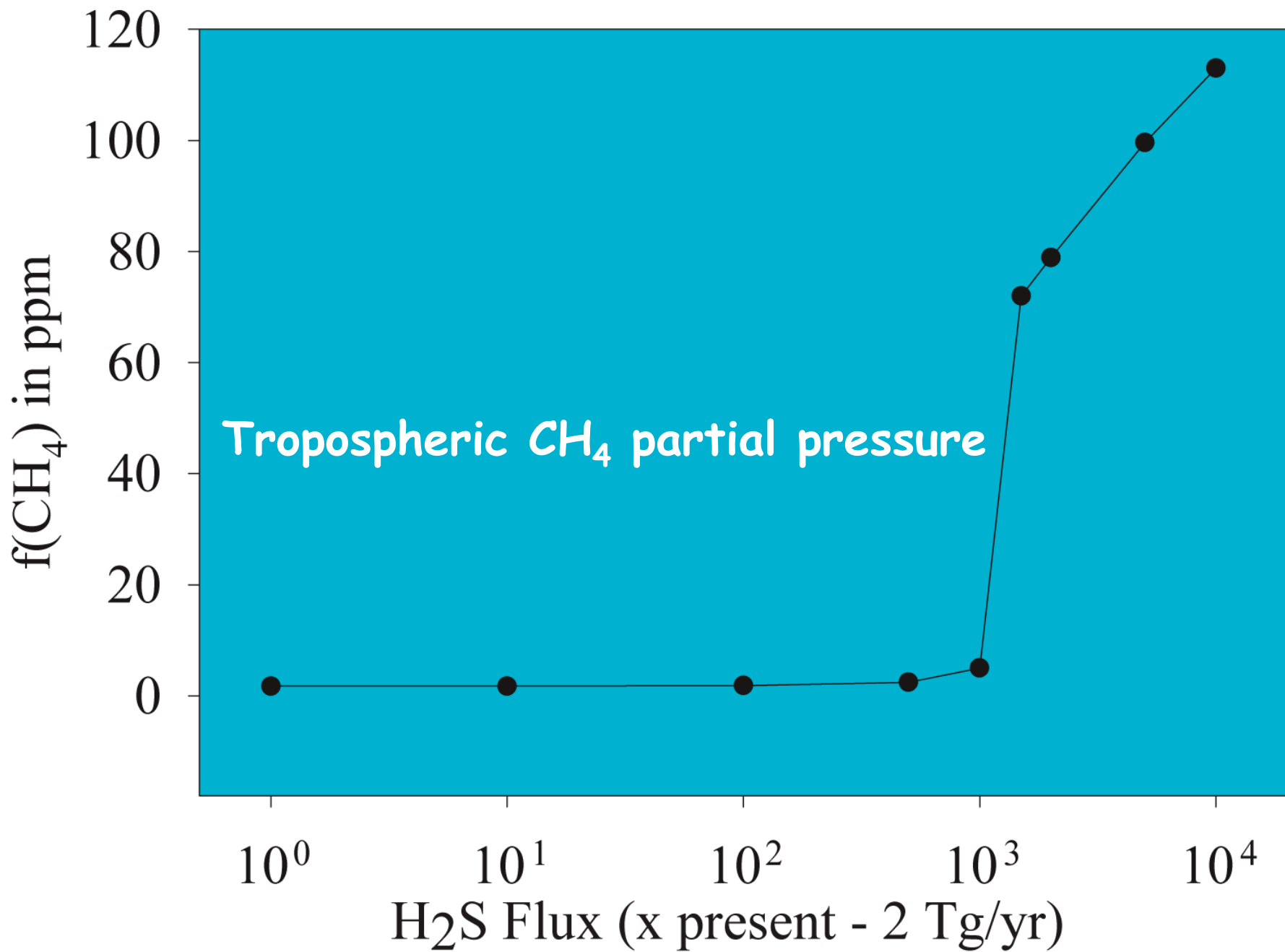
- Anoxic and sulfidic conditions



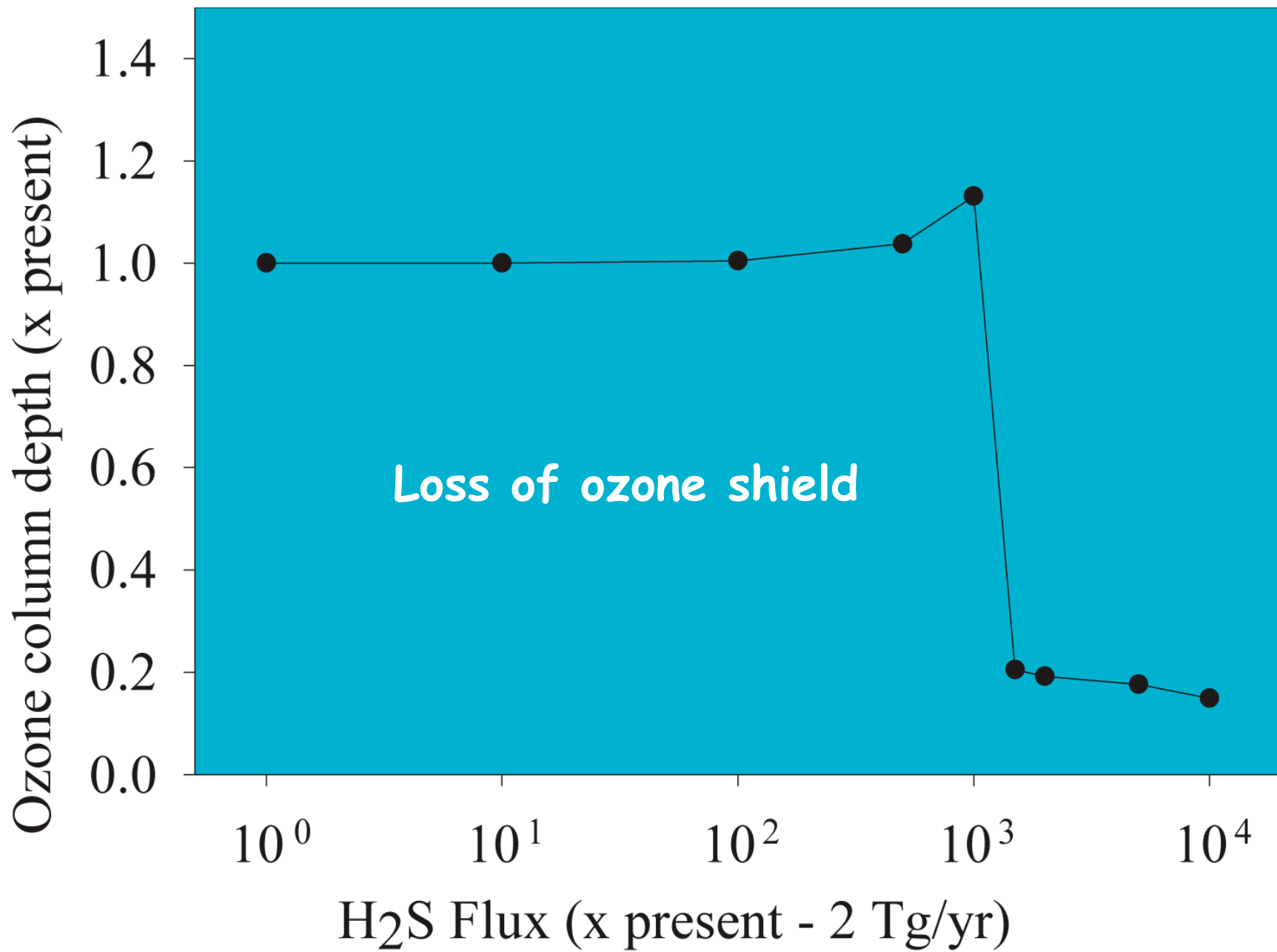














## Rotten Sulfur Brew, The Great Dying?

[www.astrobio.net/cgi-bin/h2ps.cgi?sid=672&ext=.ps](http://www.astrobio.net/cgi-bin/h2ps.cgi?sid=672&ext=.ps)

### A Welsh View: OK, Who Farted?

... "The end-Permian is puzzling," Professor Lee Kump of Penn State University told the Geological Society of America, meeting in Seattle. ...

[xo.typepad.com/blog/2003/11/ok\\_who\\_farted.html](http://xo.typepad.com/blog/2003/11/ok_who_farted.html) - 20k - Cached - Similar pages

Ananova - Stink bomb gas 'may have caused mass **extinction**'... may have taken place to the **smell** of rotten eggs ... Theories for the end-Permian **extinction** include massive volcanic ... But Dr Lee R **Kump**, professor of geosciences at ... [www.ananova.com/news/story/sm\\_835148.html](http://www.ananova.com/news/story/sm_835148.html) - 13k

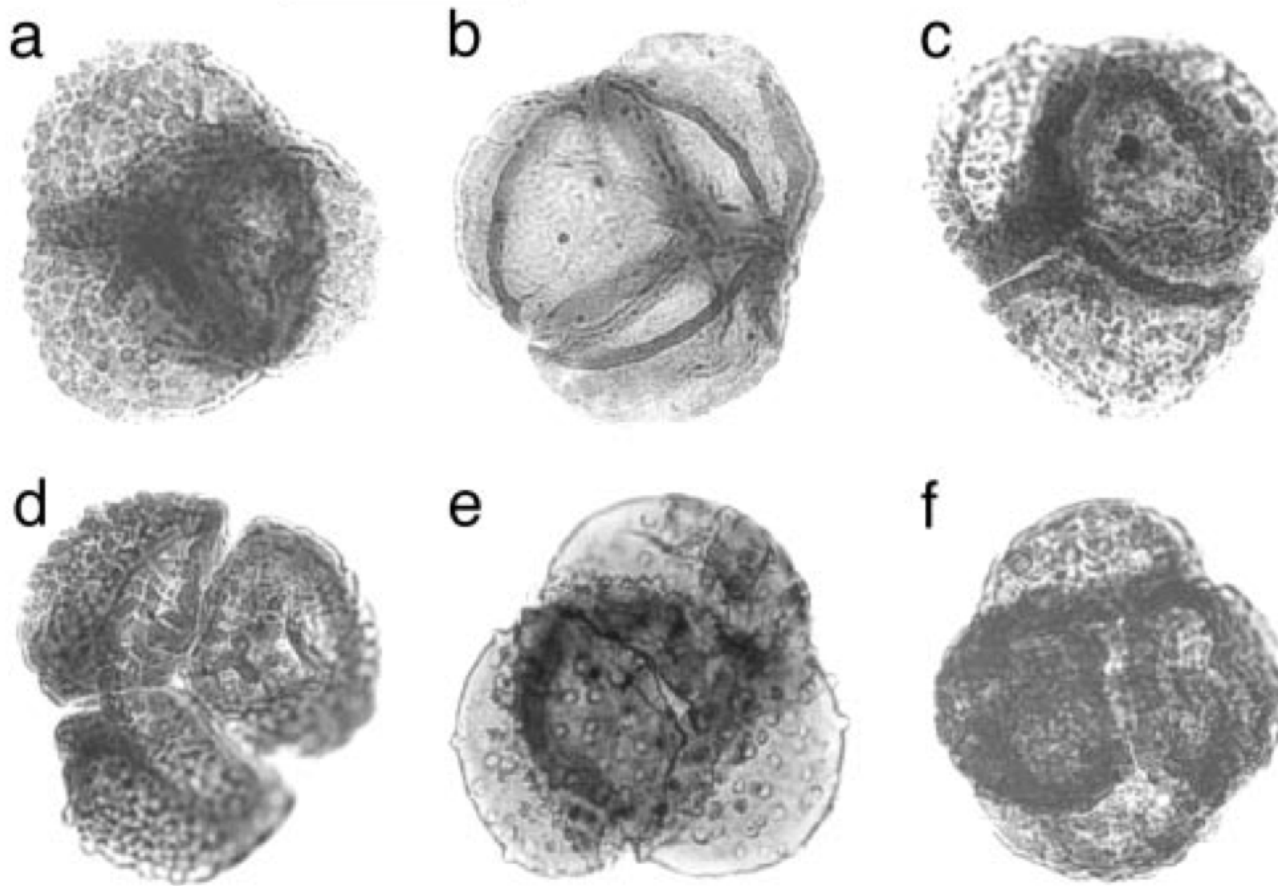
- [Cached](#) - [Similar/pages](#)

**Economist.com**

The only problem with Dr Kump's hypothesis is that he has no actual evidence for it. But he hopes to gather some soon, from rocks in Japan. And if he does, you can bet that yet another theory will come oozing down the catwalk to sneer at it.







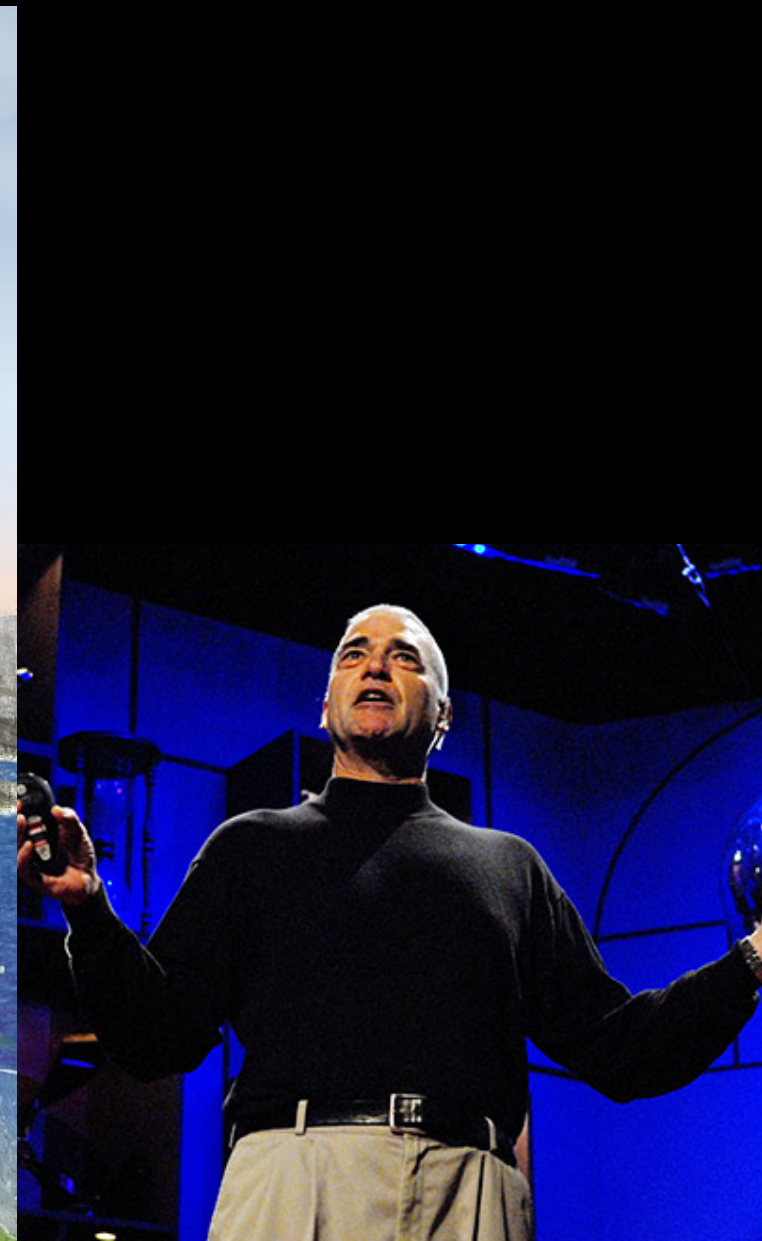
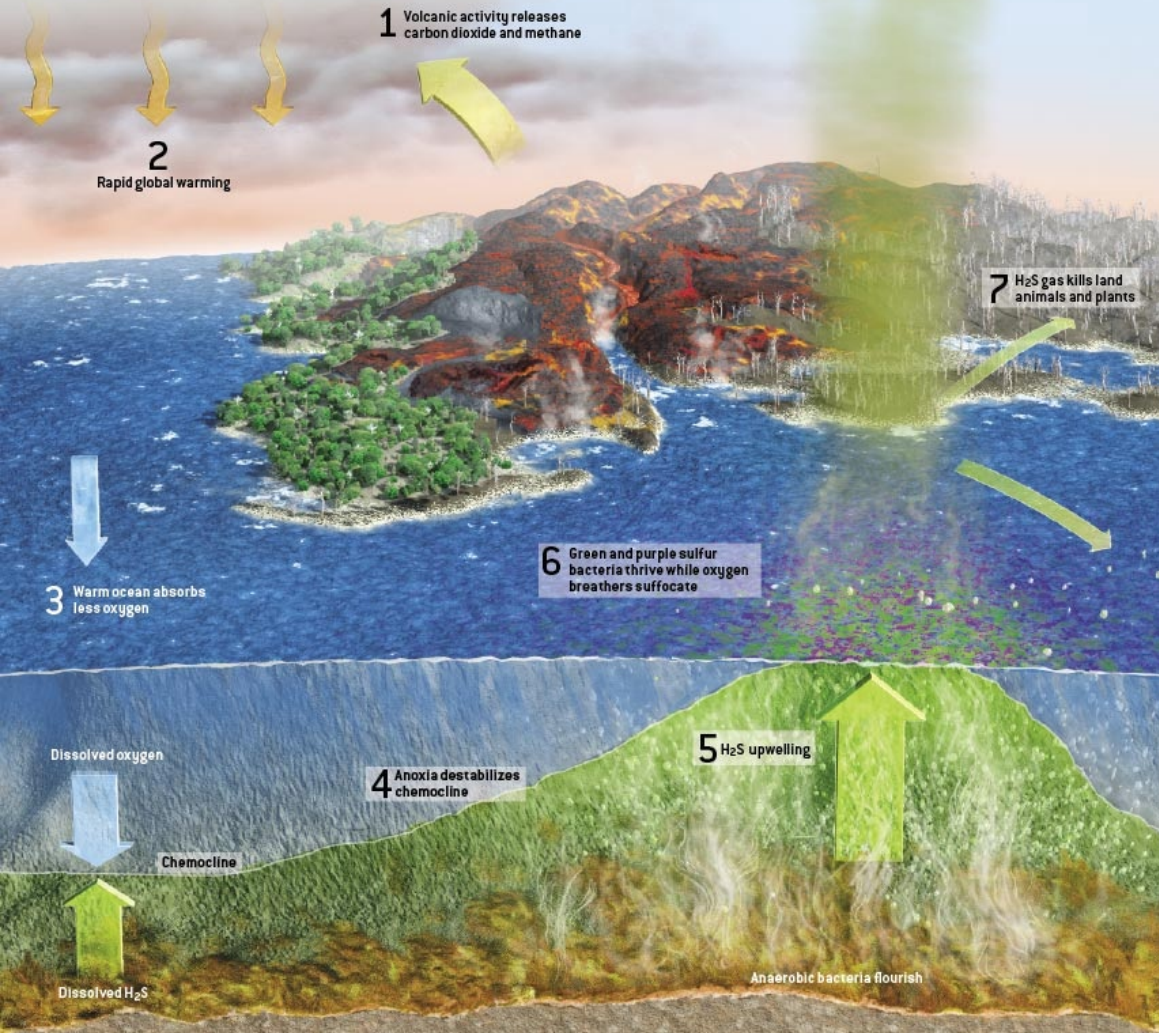
**Fig. 1.** Selection of the latest Permian microspores of heterosporous lycopids in tetrads from the Wordie Creek Formation, southern Jamesonland, East Greenland. Specimens can be assigned to various species of microspore form-genera *Lundbladispora* (a, c, e, and f), *Densoisporites* (b), and *Uvaesporites* (d). (Scale bar = 50  $\mu\text{m}$ .)

## Environmental mutagenesis during the end-Permian ecological crisis

Henk Visscher<sup>\*†</sup>, Cindy V. Looy<sup>\*‡</sup>, Margaret E. Collinson<sup>§</sup>, Henk Brinkhuis<sup>\*</sup>,  
Johanna H. A. van Konijnenburg-van Cittert<sup>\*¶</sup>, Wolfram M. Kürschner<sup>\*</sup>, and Mark A. Sephton<sup>||</sup>

# KILLER GREENHOUSE EFFECT

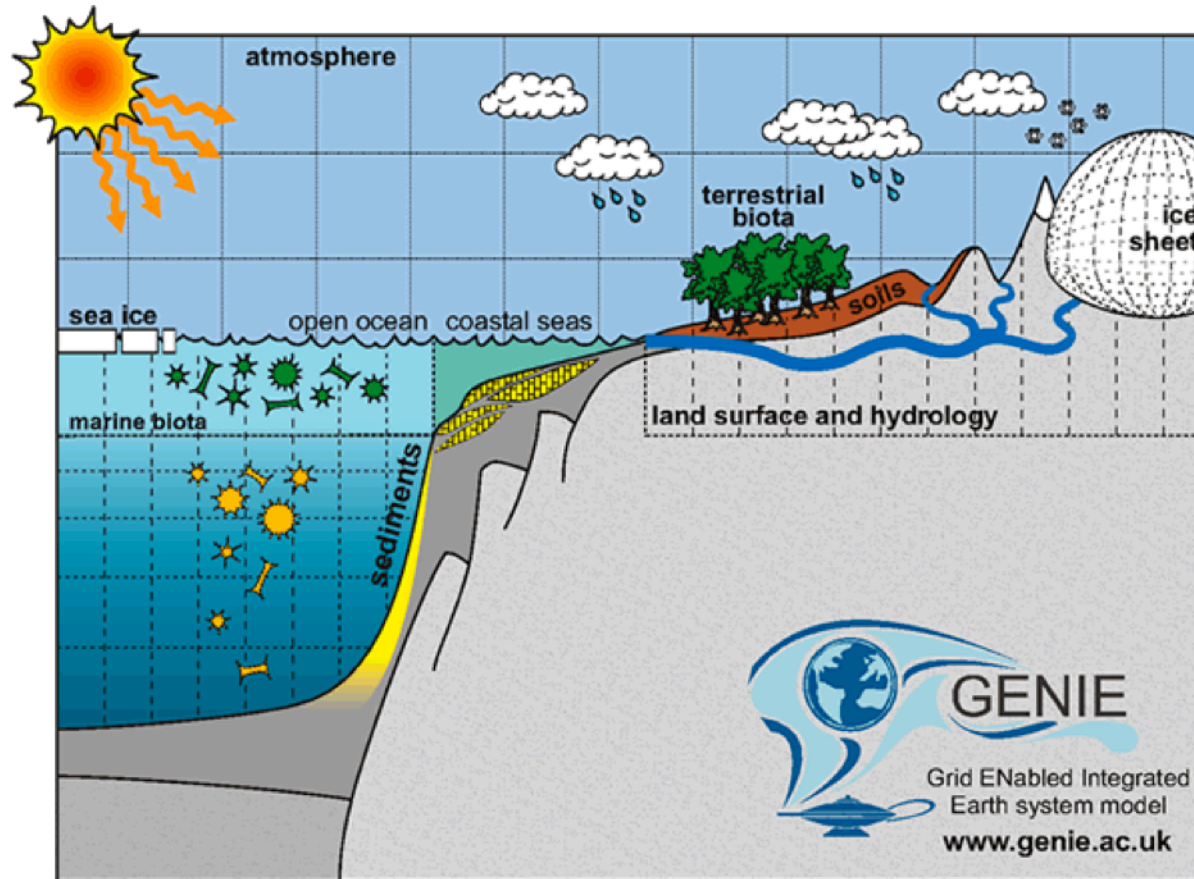
A new model for mass extinctions at the end of the Permian period 251 million years ago and the end Triassic 50 million years later explains how intense global warming could trigger deaths in the sea and on land. Trouble begins with widespread volcanic activity that releases enormous volumes of carbon dioxide and methane [1]. The gases cause rapid global warming [2]. A warmer ocean absorbs less oxygen from the atmosphere [3]. Low oxygen (anoxia) destabilizes the chemocline, where oxygenated water meets water permeated with hydrogen sulfide ( $H_2S$ ) generated by bottom-dwelling anaerobic bacteria [4]. As  $H_2S$  concentrations build and oxygen falls, the chemocline rises abruptly to the ocean surface [5]. Green and purple photosynthesizing sulfur bacteria, which consume  $H_2S$  and normally live at chemocline depth, now inhabit the  $H_2S$ -rich surface waters while oxygen-breathing ocean life suffocates [6].  $H_2S$  also diffuses into the air, killing animals and plants on land [7] and rising to the troposphere to attack the planet's ozone layer [8]. Without the ozone shield, the sun's ultraviolet (UV) radiation kills remaining life [9].



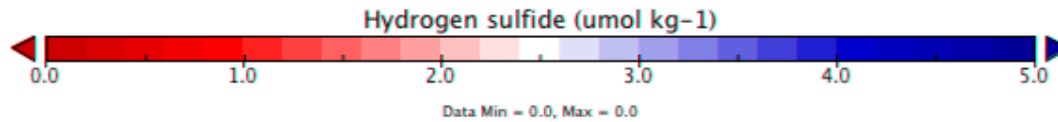
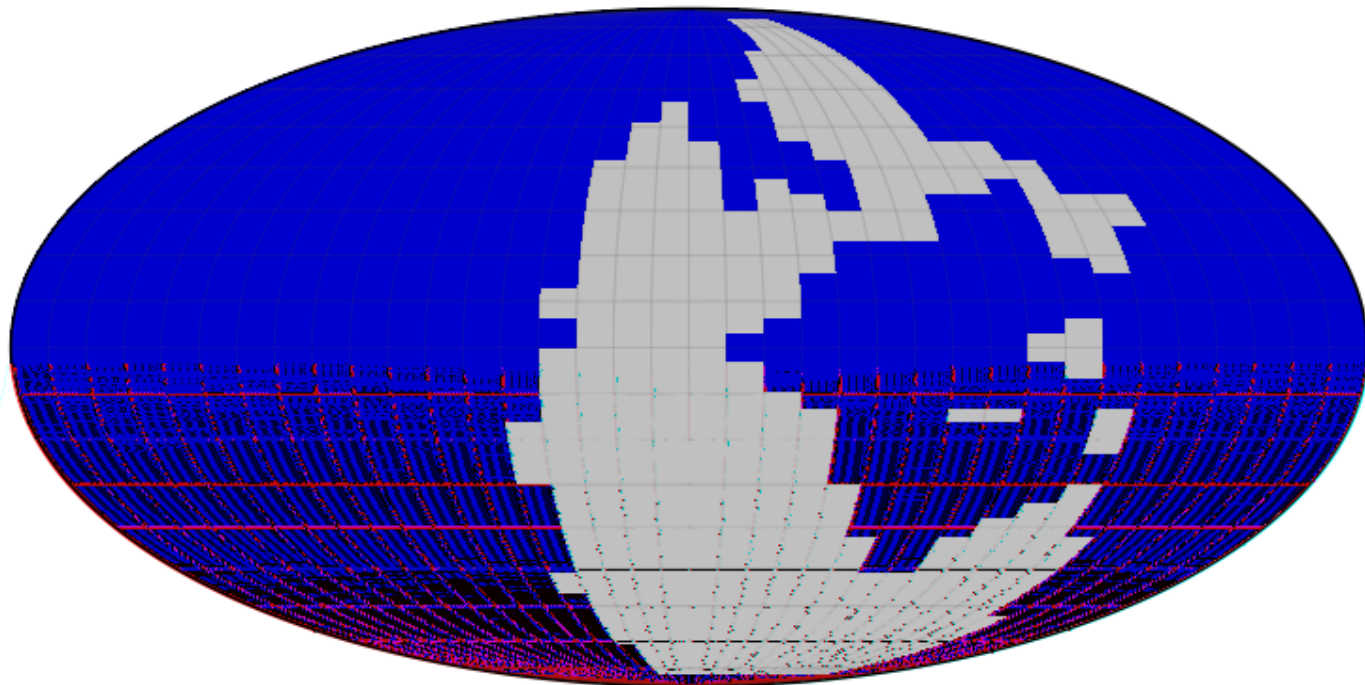
P. Ward (Sci Am. 2006)



- Spatially-resolved Earth system models

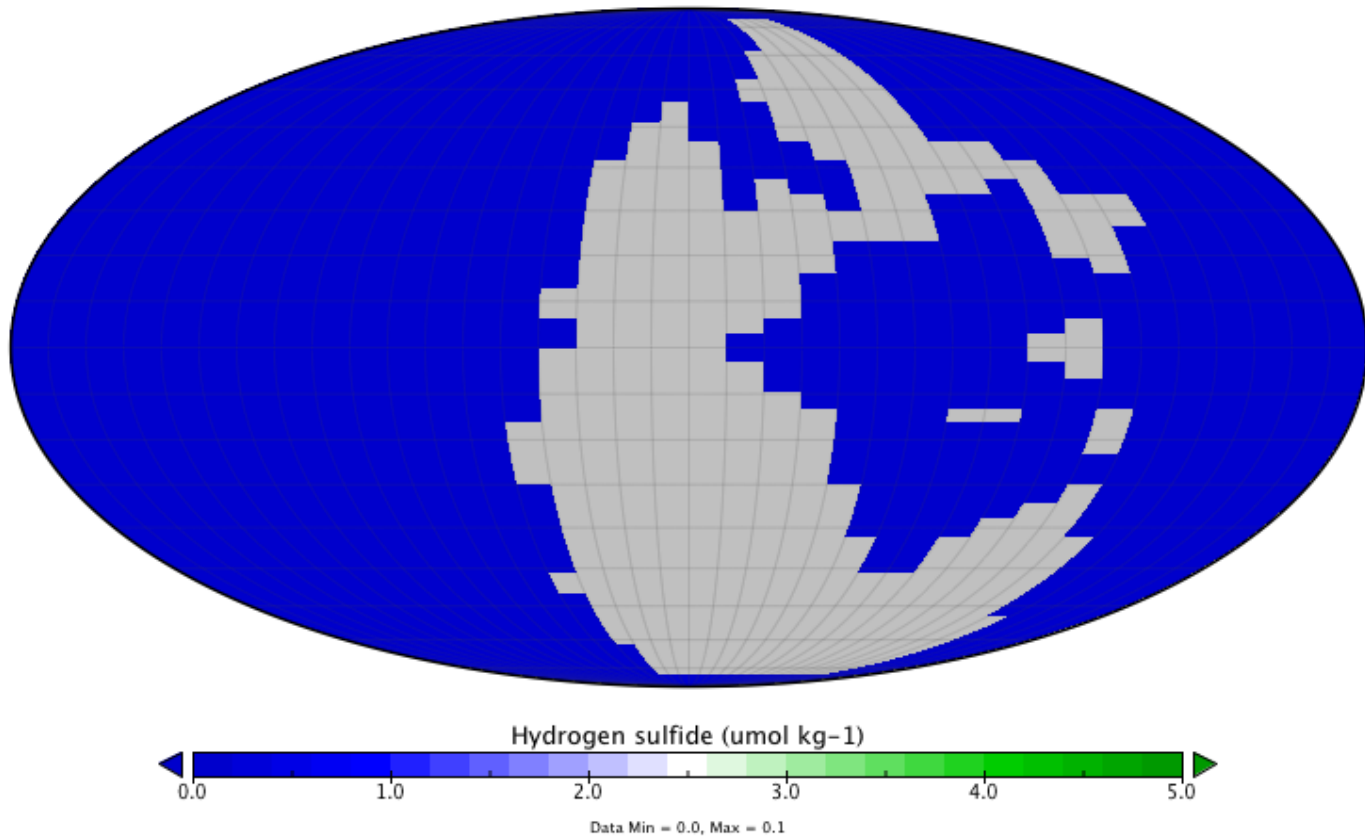


# SURFACE OCEAN $\text{H}_2\text{S}$



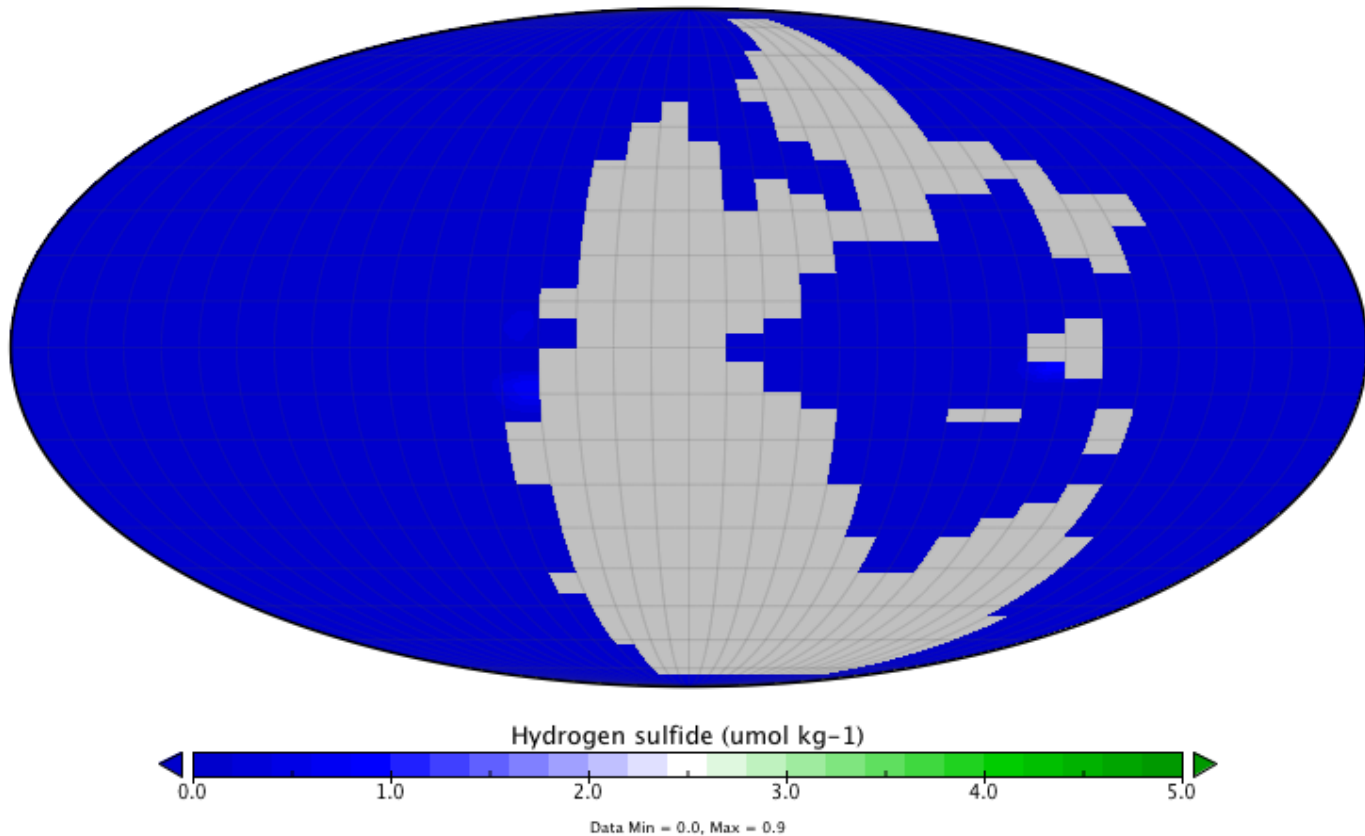
$1 \times \text{PO}_4^{3-}$

# SURFACE OCEAN $\text{H}_2\text{S}$



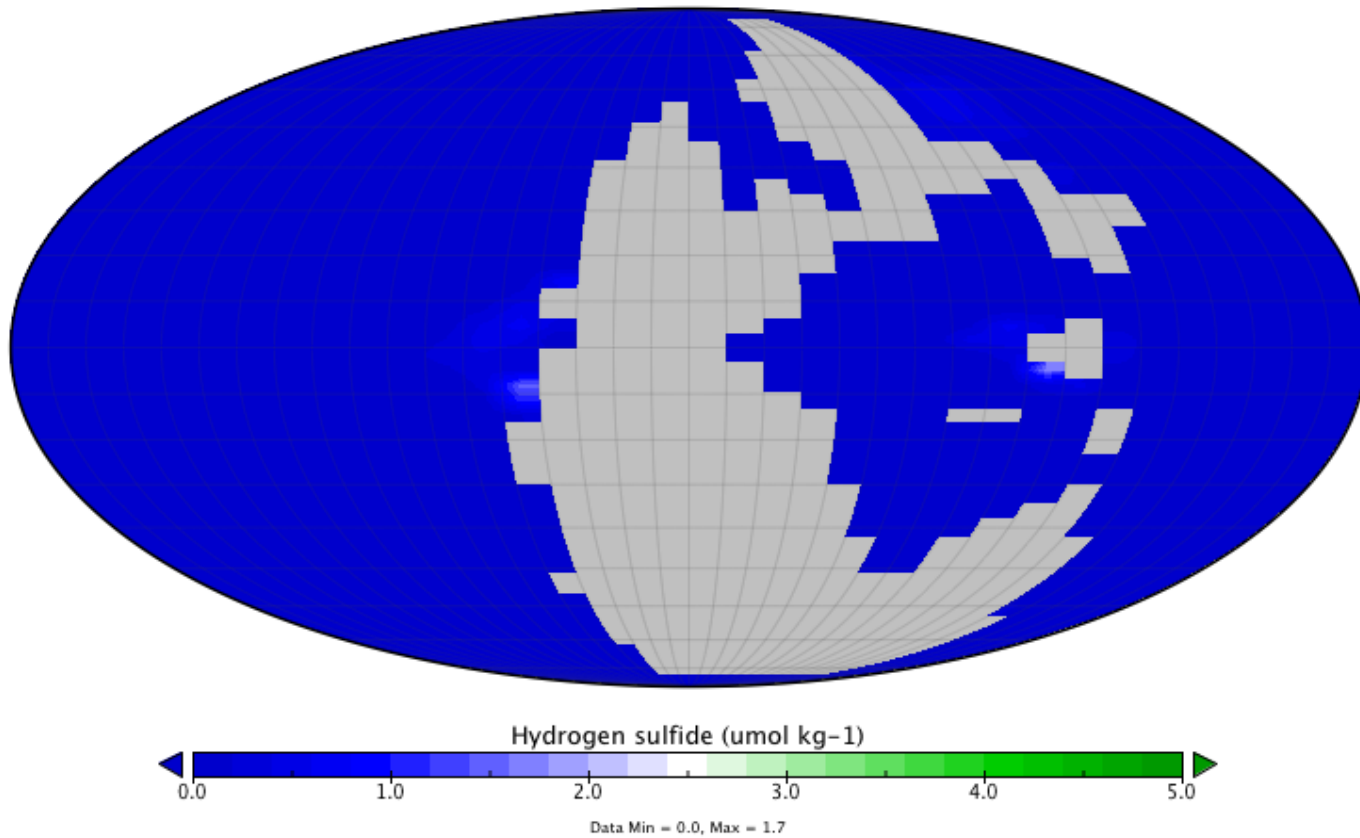
$2x \text{PO}_4^{3-}$

# SURFACE OCEAN $\text{H}_2\text{S}$



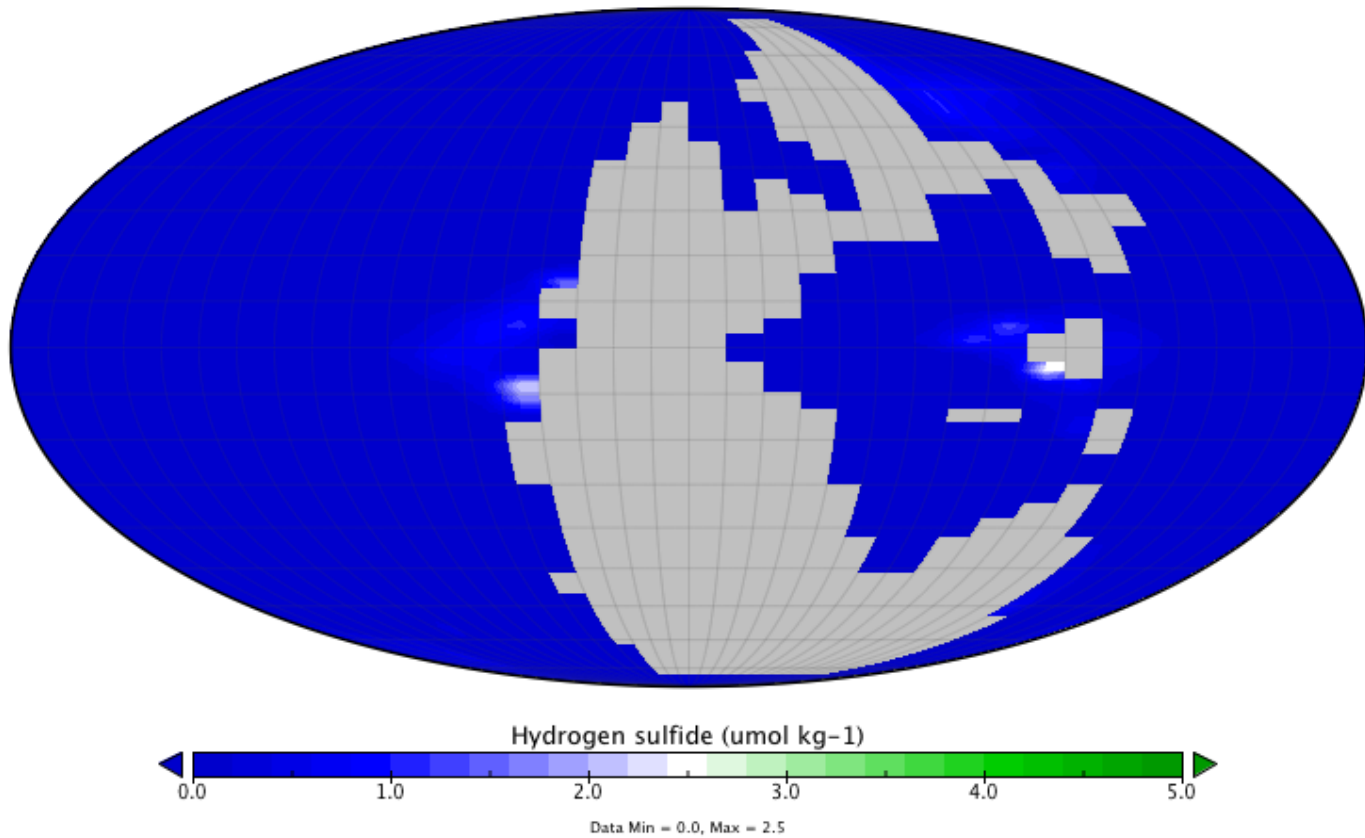


# SURFACE OCEAN $\text{H}_2\text{S}$

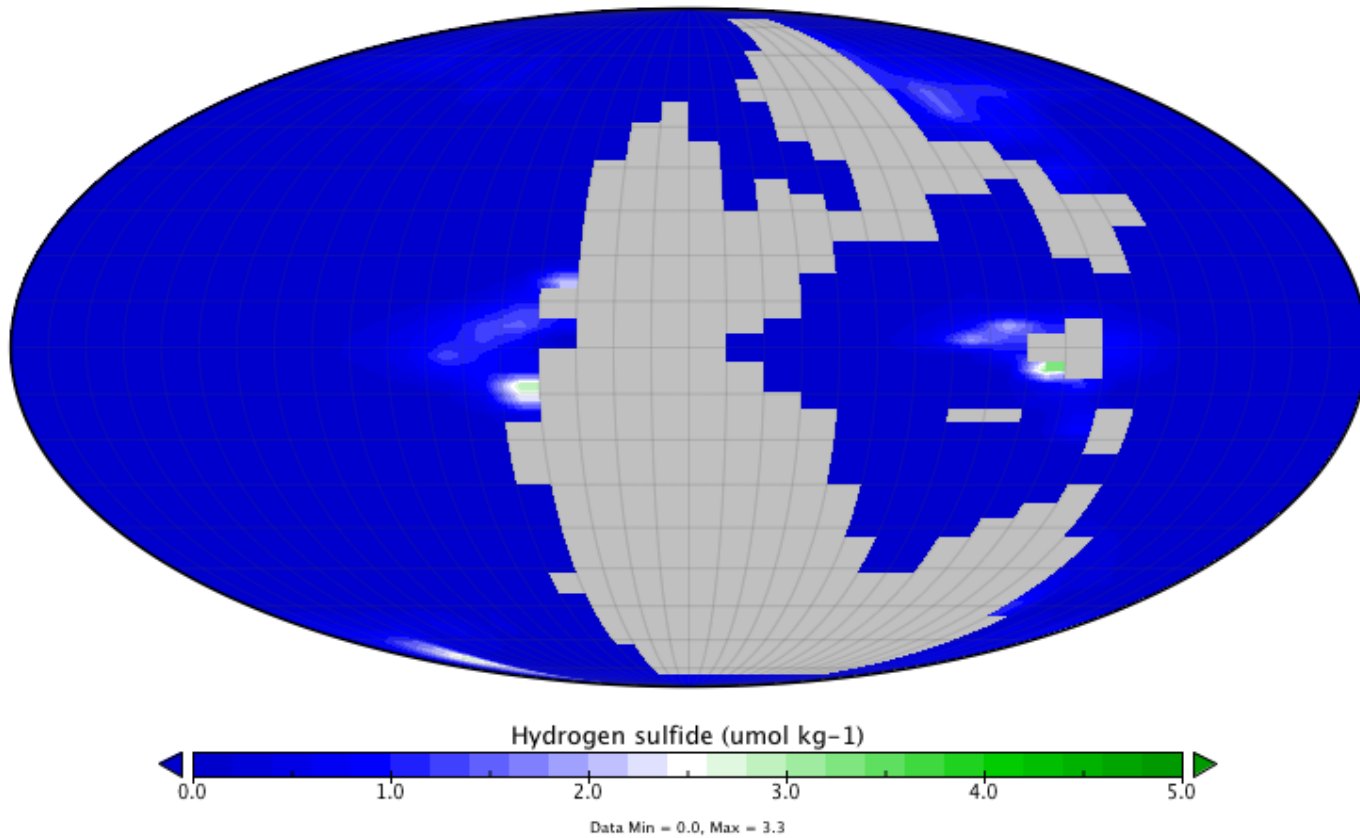


4x  $\text{PO}_4^{3-}$

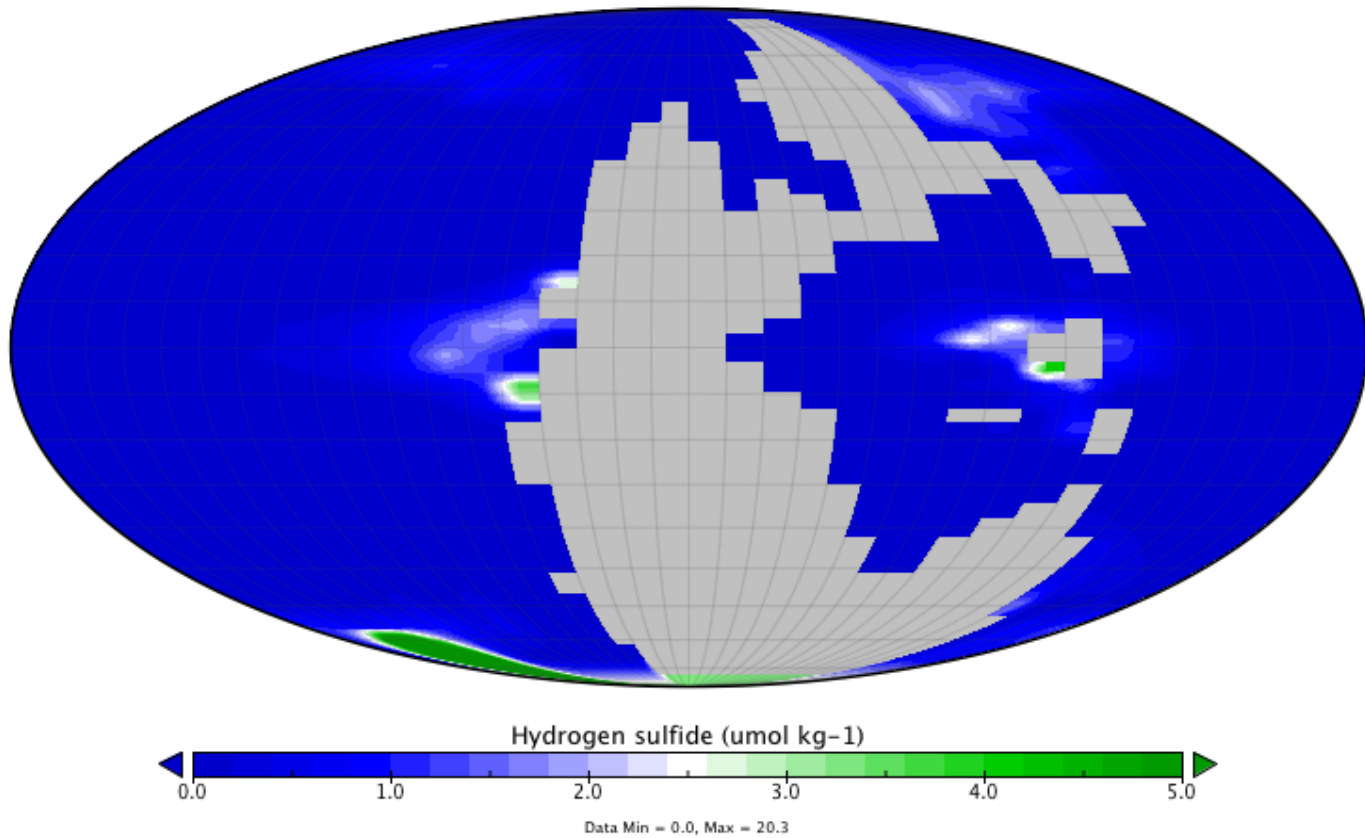
# SURFACE OCEAN H<sub>2</sub>S



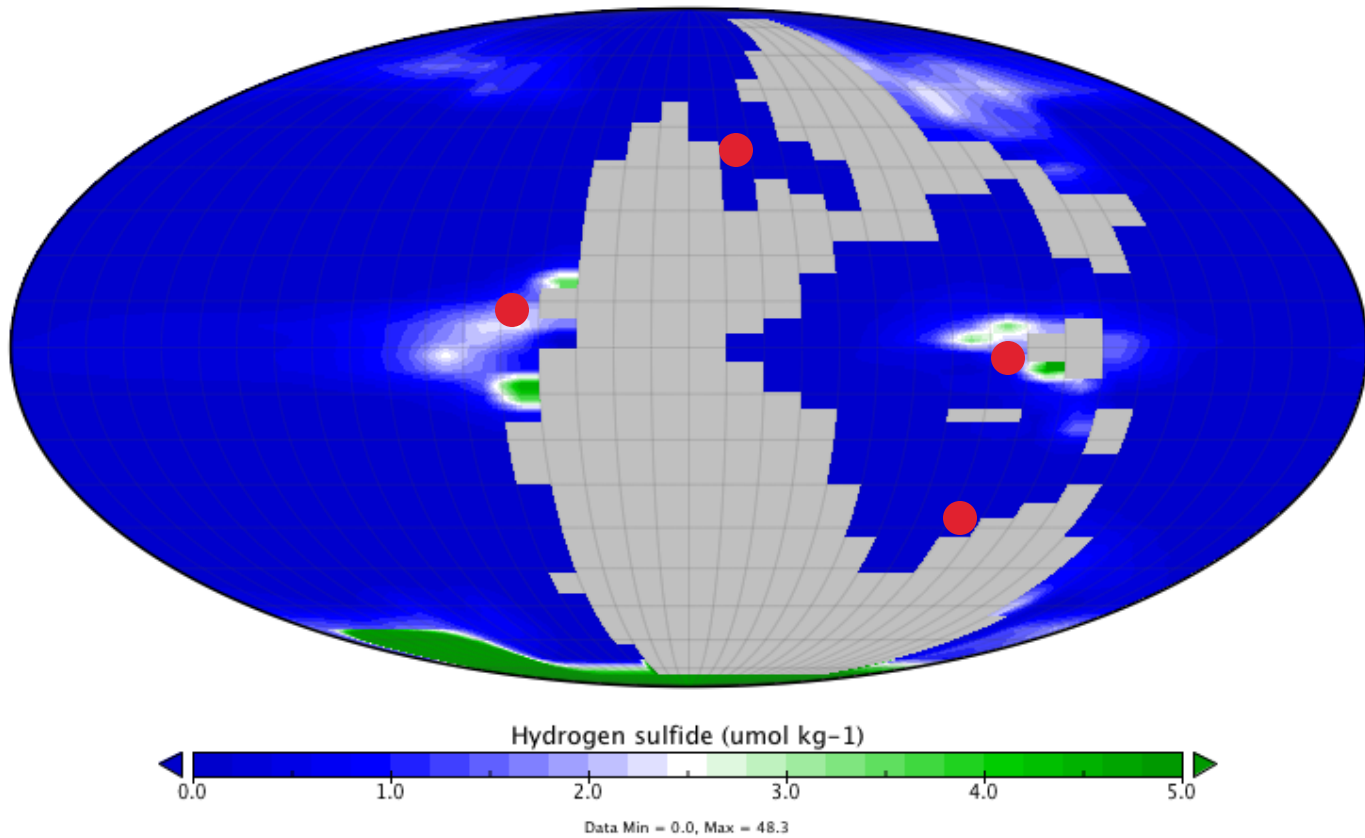
# SURFACE OCEAN $\text{H}_2\text{S}$



# SURFACE OCEAN H<sub>2</sub>S

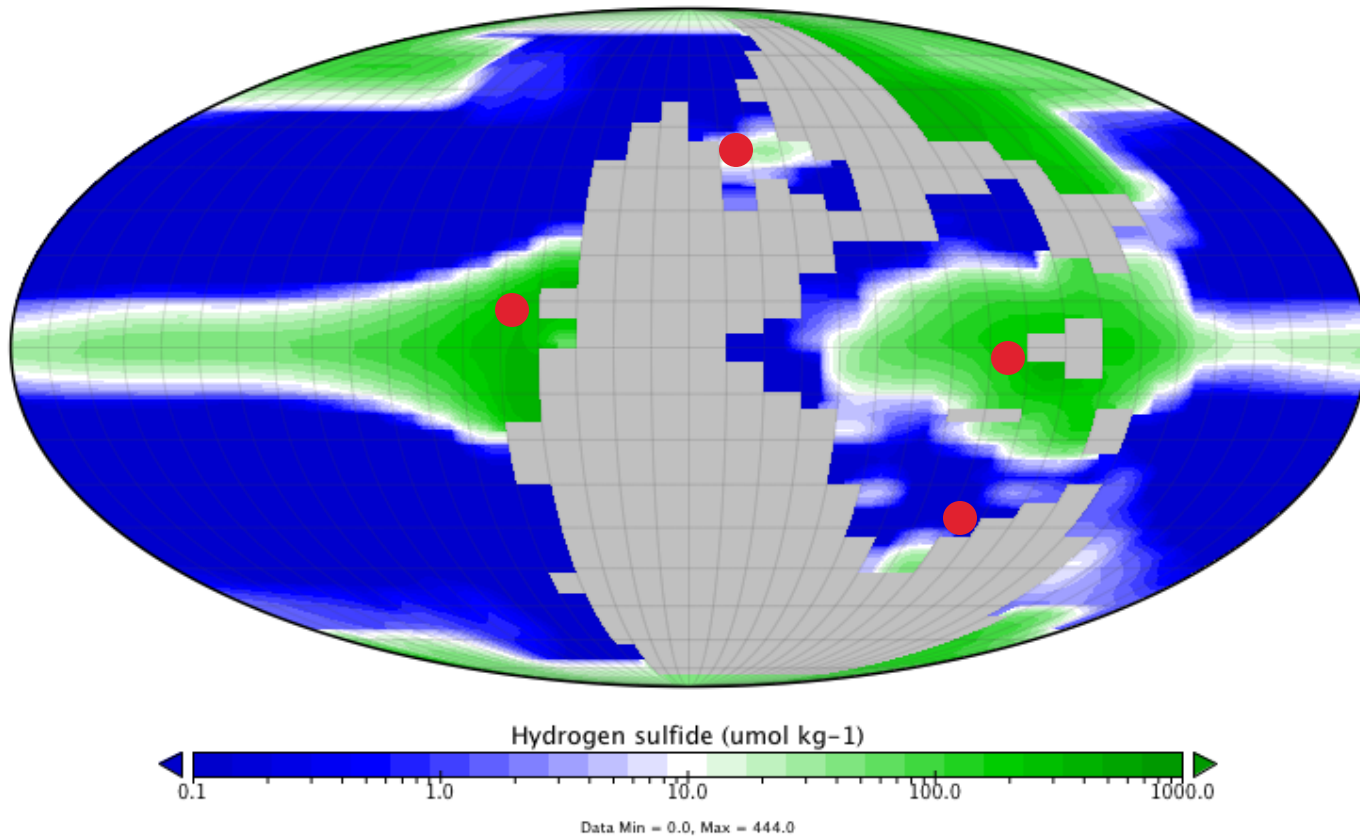


# SURFACE OCEAN H<sub>2</sub>S





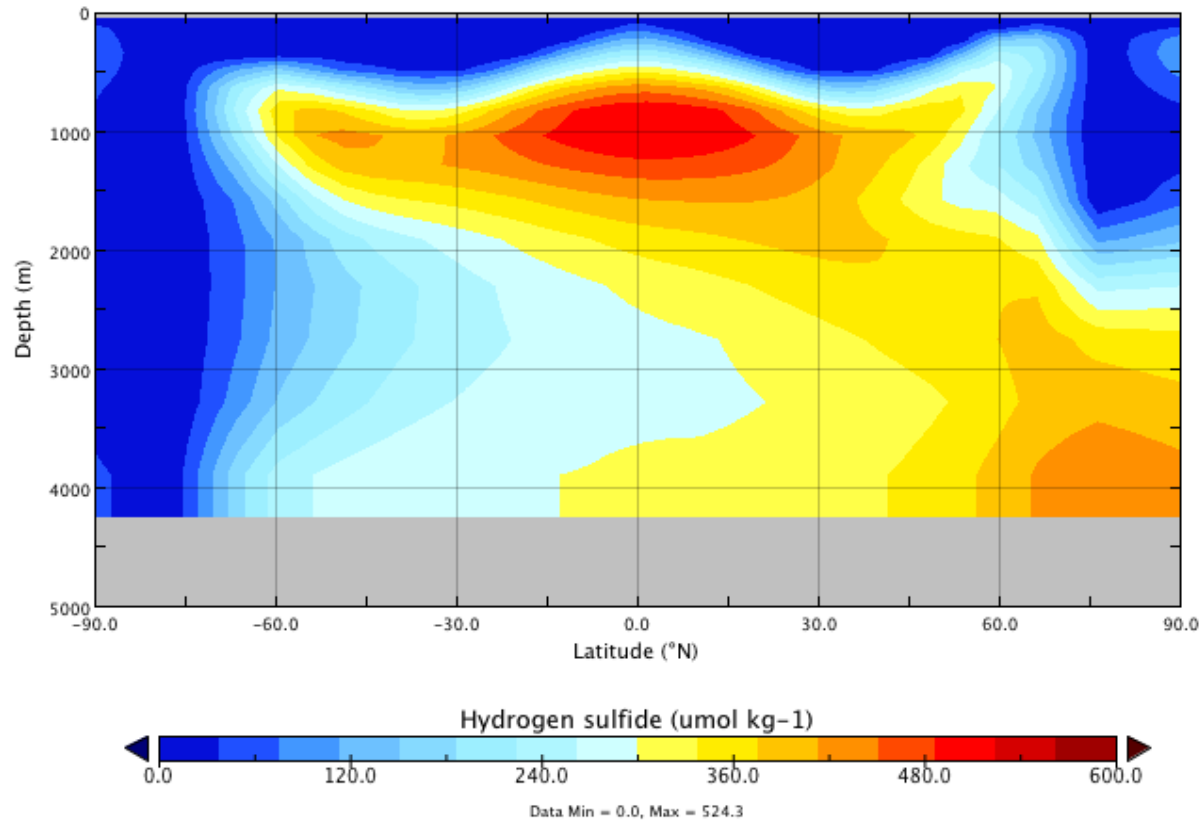
# SURFACE OCEAN $H_2S$



80-175 M DEPTH

8X  $\text{PO}_4^{3-}$

# HYDROGEN SULFIDE

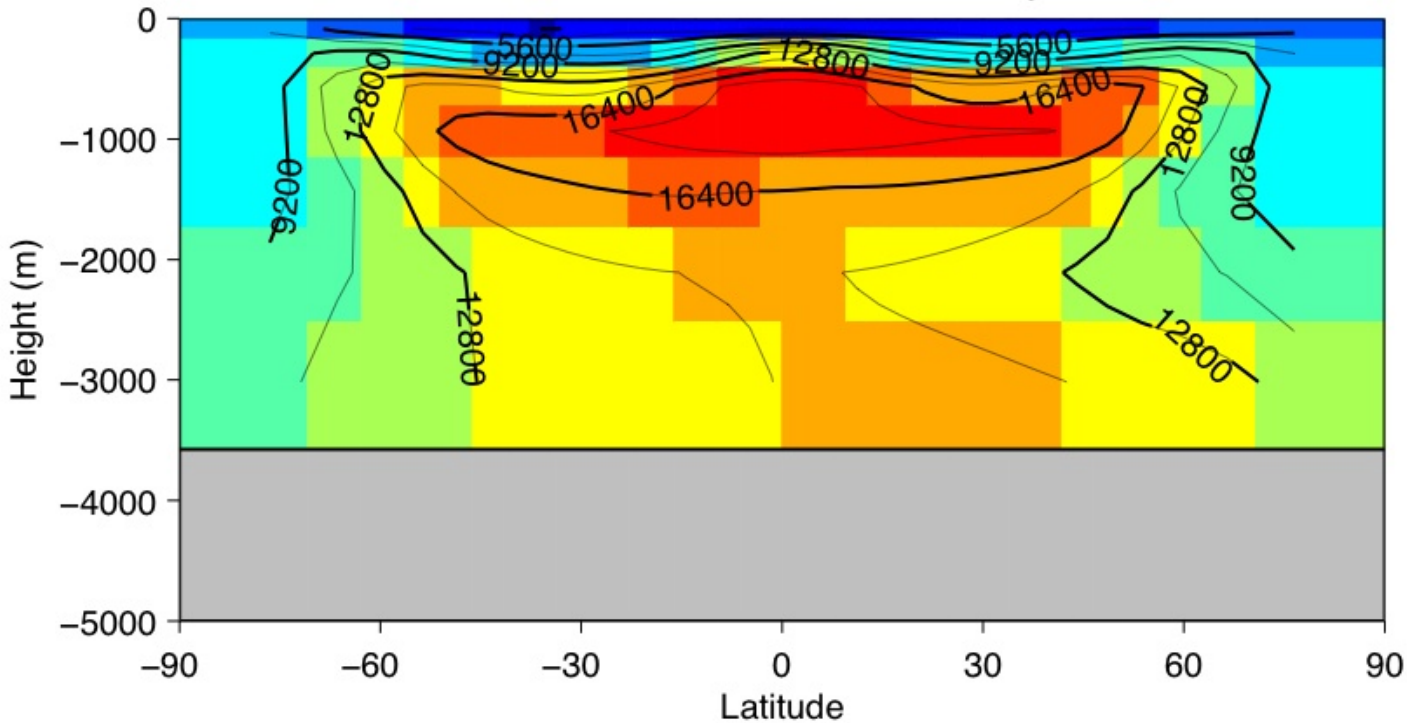


PANTHALASSIC OCEAN

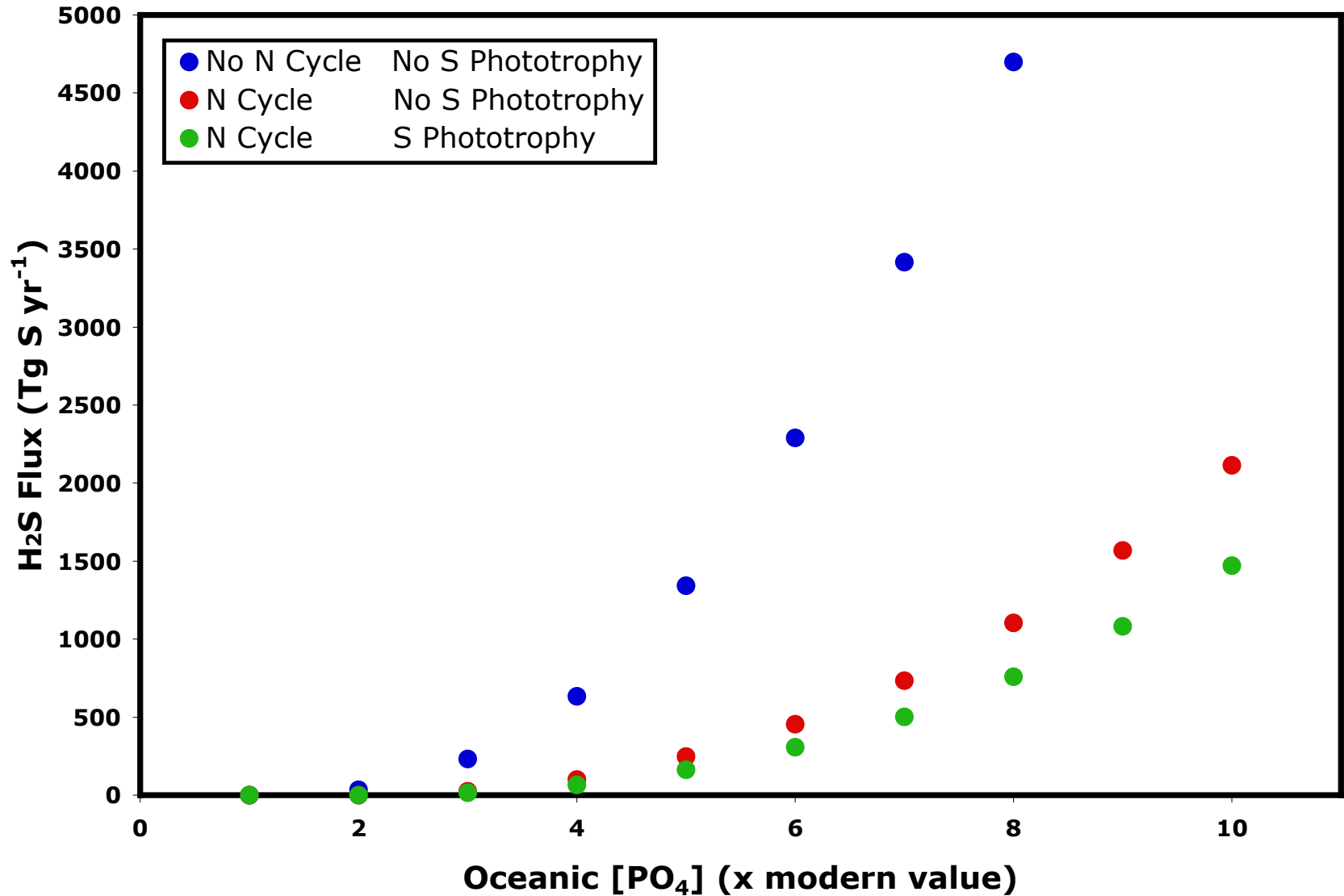
8X PO<sub>4</sub><sup>3-</sup>

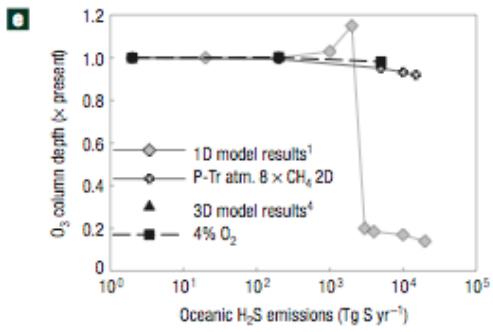
Run: 060209b / Time: 60000 / Data: carb fug CO2 / i = 5

ppm CO<sub>2</sub>

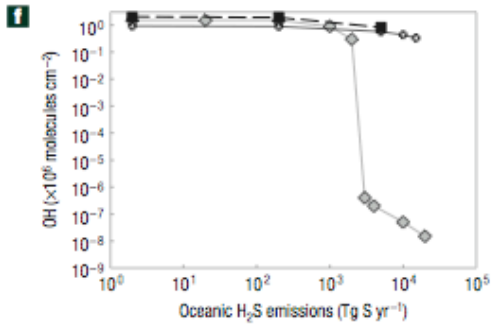


## More realistic models have reduced fluxes

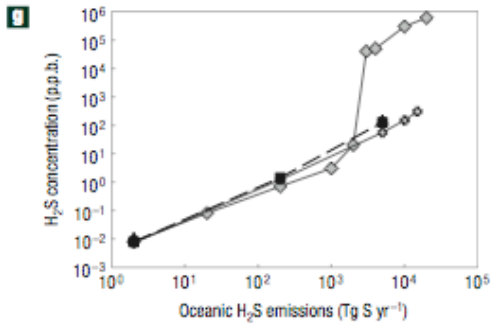




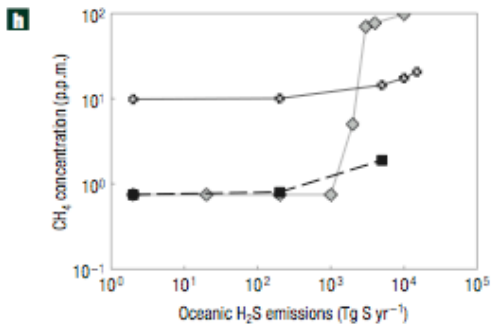
$O_3$



OH



$H_2S$

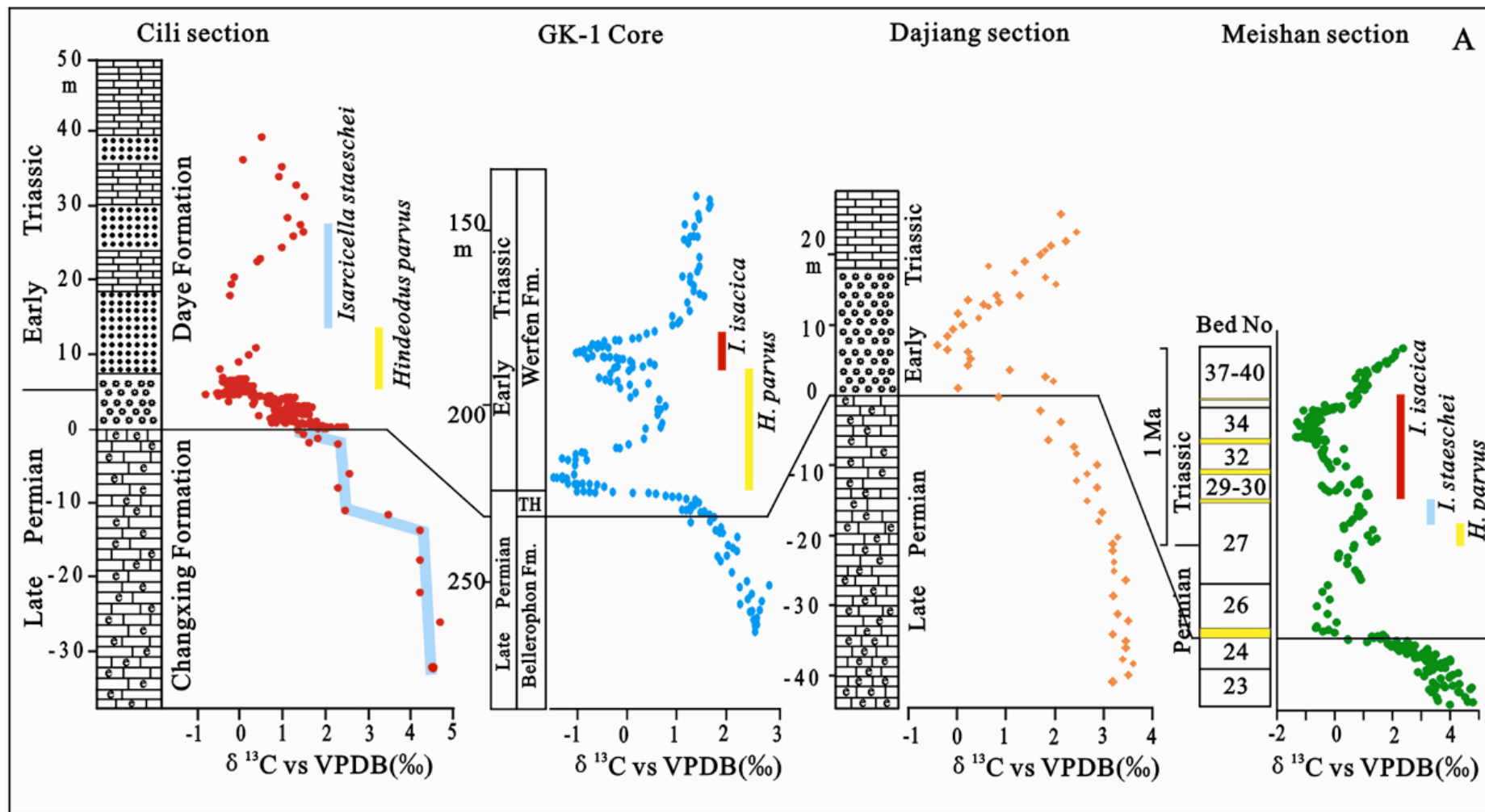


$CH_4$



# Is the H<sub>2</sub>S Hypothesis “Dead”?

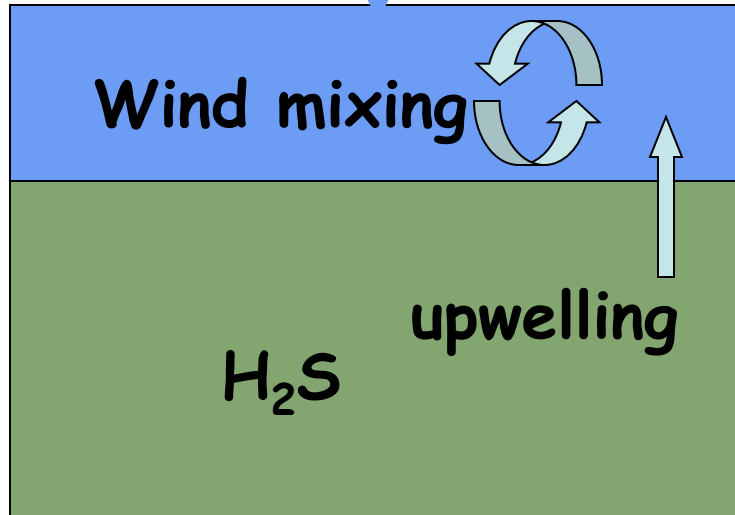
- More sophisticated ocean models produce less H<sub>2</sub>S in surface waters (but PZE still widespread)
- More sophisticated atmospheric models yield much lower H<sub>2</sub>S values and no ozone depletion at similar fluxes to earlier work (chronic rather than acute stress on biota)



**Keeping euxinia at depth:** is the Black Sea the appropriate model for an euxinic ocean?

**Atmospheric  $O_2$**

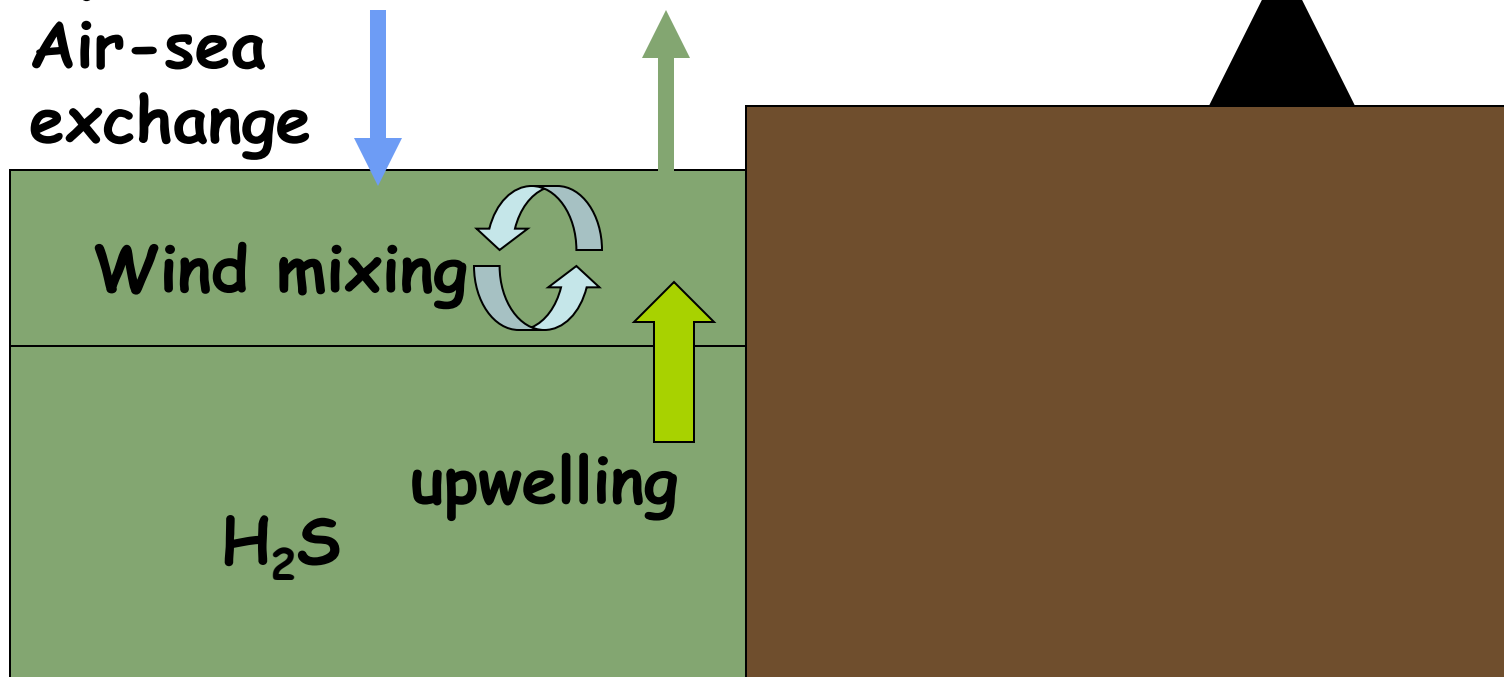
Air-sea  
exchange



Wind-mixed layer oxygenated today (e.g., in Black Sea) because supply of  $O_2$  across air-sea interface exceeds upwelling of  $H_2S$

Upwelling flux of  $\text{H}_2\text{S}$  comes to exceed in-mixing of  $\text{O}_2$

**Atmospheric  $\text{O}_2$**



Because of increasing  $[\text{H}_2\text{S}]$  and/or decreasing atmospheric  $p\text{O}_2$